

How much is global gross primary production
of the terrestrial biosphere?

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Global terrestrial GPP

- Gross primary production, GPP, is the largest flux in terrestrial carbon cycle, can not be measured directly.
- Current estimates vary widely:

By modeling. Beer et al. (2010):107– 139Gt C year⁻¹

From NPP. Piao et al. (2010):100– 157 Gt C year⁻¹

CO¹⁶O¹⁸. Welp et al.(2011):150– 175 Gt C year⁻¹

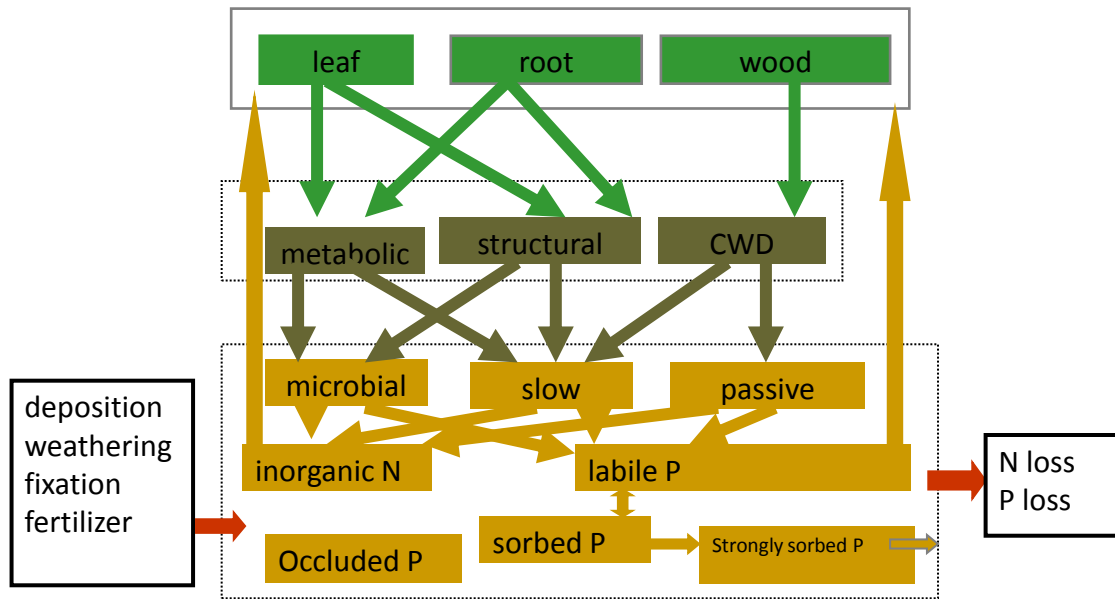
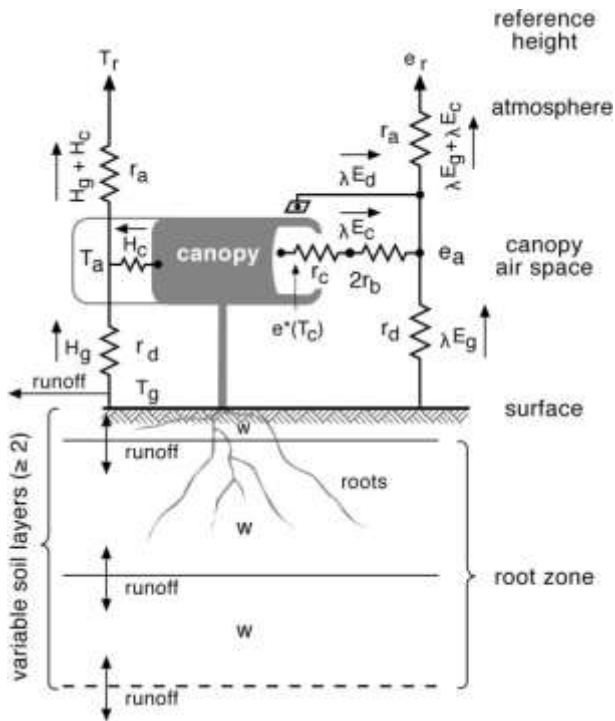
Issues with various estimates

- Modeling
 - accuracy depends on calibration data
 - Observations are very sparse for some regions
 - Issues with nighttime flux measurements
- NPP/GPP ratio
 - Depending how good estimates of NPP are
 - the ratio quite variable, quite low under P limiting
- Isotope method
 - some unverified assumptions.
- Global C4 fraction may be underestimated.

Another modeling approach

- water and carbon are strongly coupled, and much of our present knowledge is captured in the Australian community land surface model (CABLE).
- Two questions:
 - How can we use CABLE to predict global GPP subject to the constraints of the “observed” global latent heat fluxes, global LAI and prior information about CABLE parameters?
 - What is the uncertainty of the predicted GPP?

The CABLE model



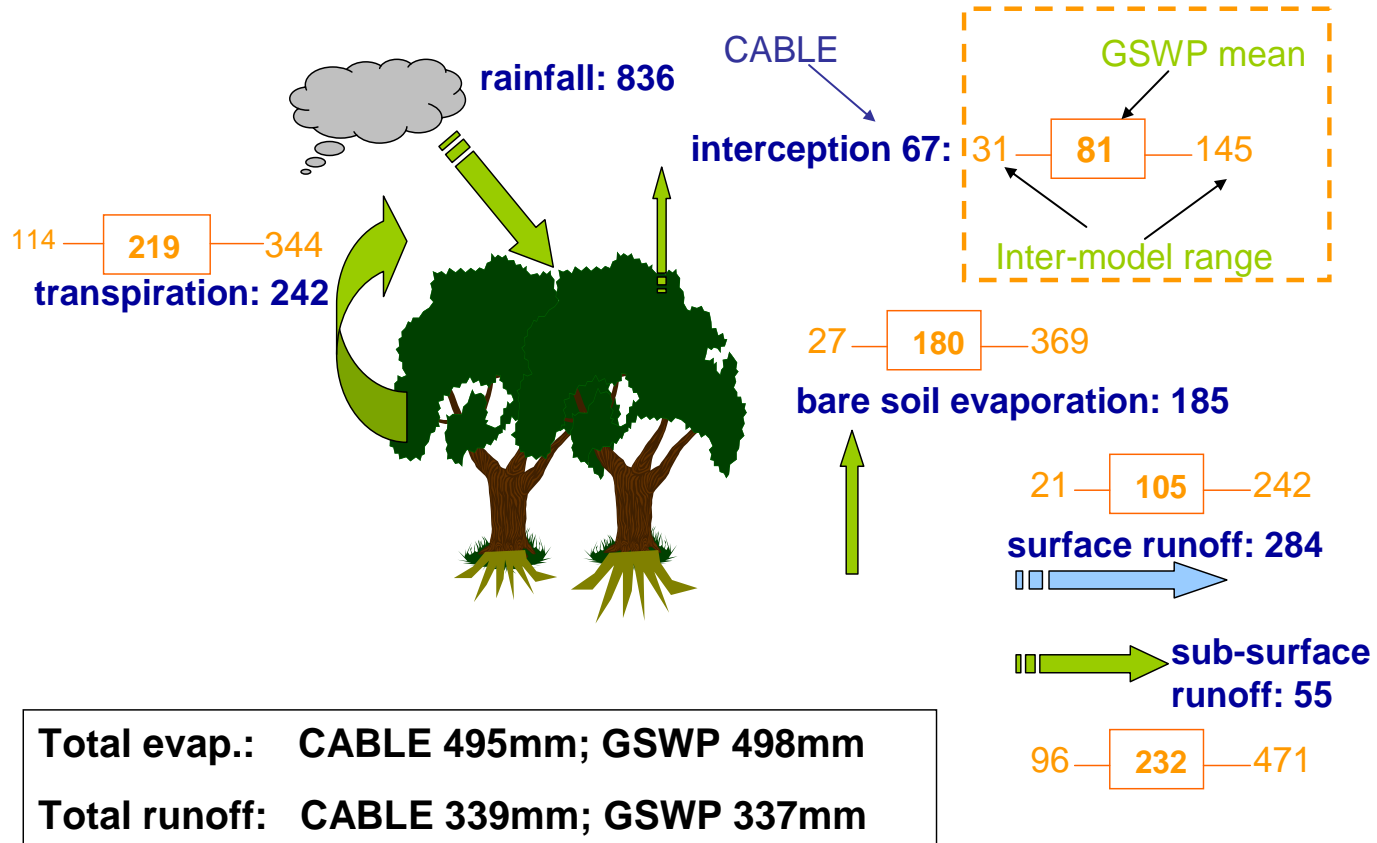
Biophysics



Biogeochemistry (C, N, P) = CABLE

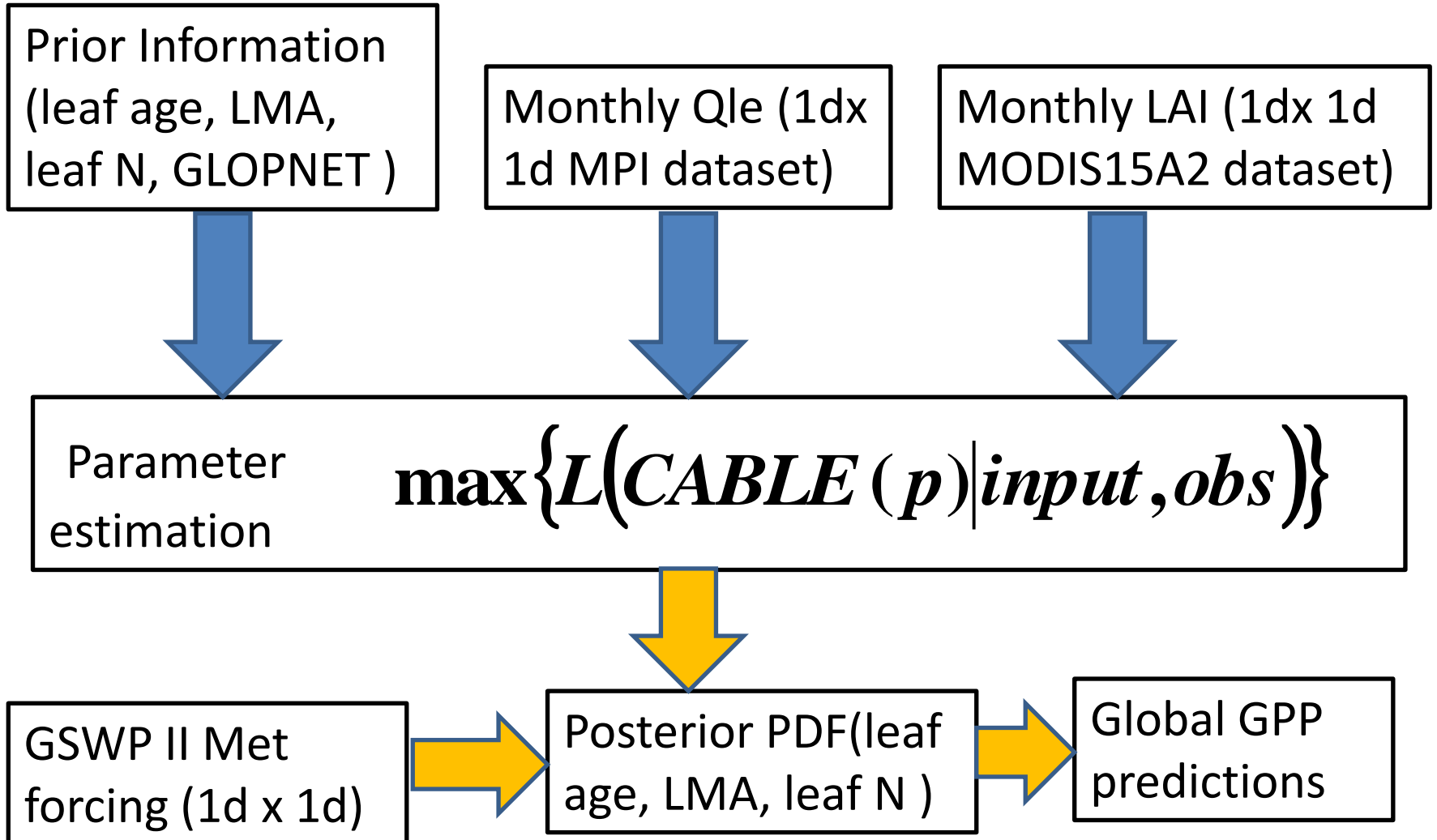
Global water budget by default CABLE

Global water budget (mm yr⁻¹) against GSWP 13-model climatology
(shown in orange colour with mean and inter-model range)



From: Zhang et al. 2012

A Bayesian inversion



Why those three leaf parameters?

- H₂O and C fluxes depends on two variables:
 - **Canopy LAI**: total area of water and C fluxes
$$\text{LAI} = f_1(\text{leaf age}, \text{leaf mass per area}, \text{NPP}, a_{\text{leaf}})$$
 - **Leaf nitrogen**: affecting the rate of water/carbon fluxes per unit LAI.
$$V_{\text{cmax}} = f_2(\text{leaf N}\%)$$
- Three parameters are optimized using PEST
- why three? (only learnt new PEST last week)

The uncertainty of model predictions (linear theory)

$$\hat{F}_{gpp} = y^T \hat{p}$$

$$\sigma_{GPP}^2 = \underbrace{y^T (I - R)C(p)(I - R)^T y}_{\text{model structure error}} + \underbrace{y^T GC(\varepsilon)G^T y}_{\text{model-obs mismatch}}$$

R: model resolution matrix

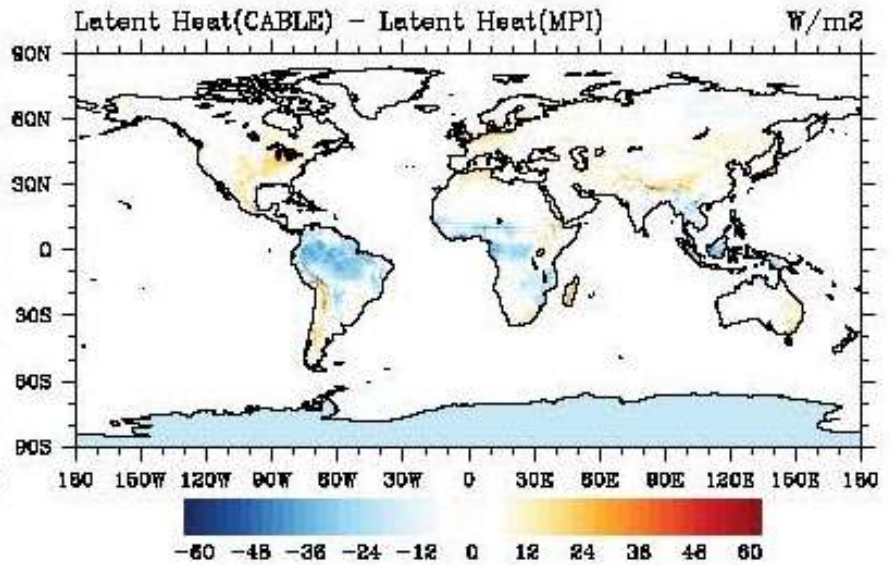
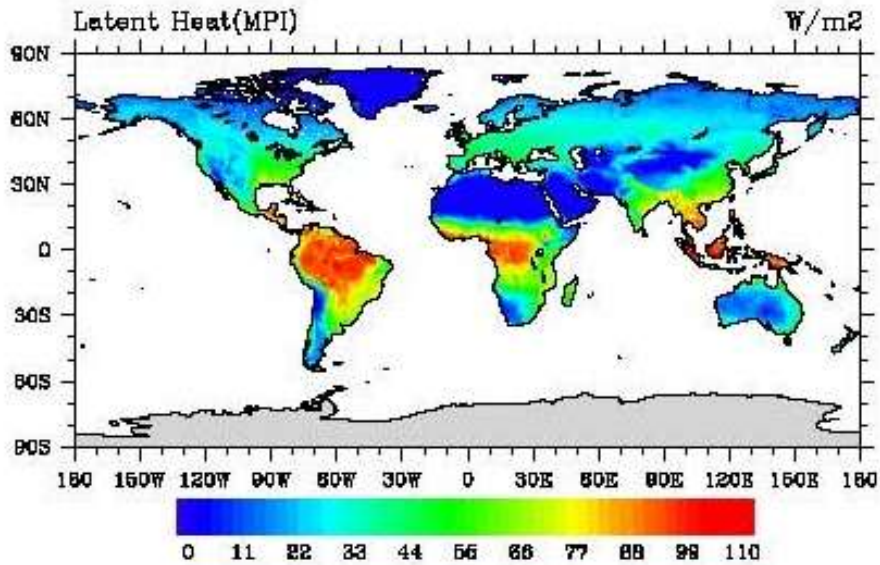
I: identity matrix

C(p): parameter covariance matrix

G: parameter solution matrix

C(ε): measurement error covariance matrix

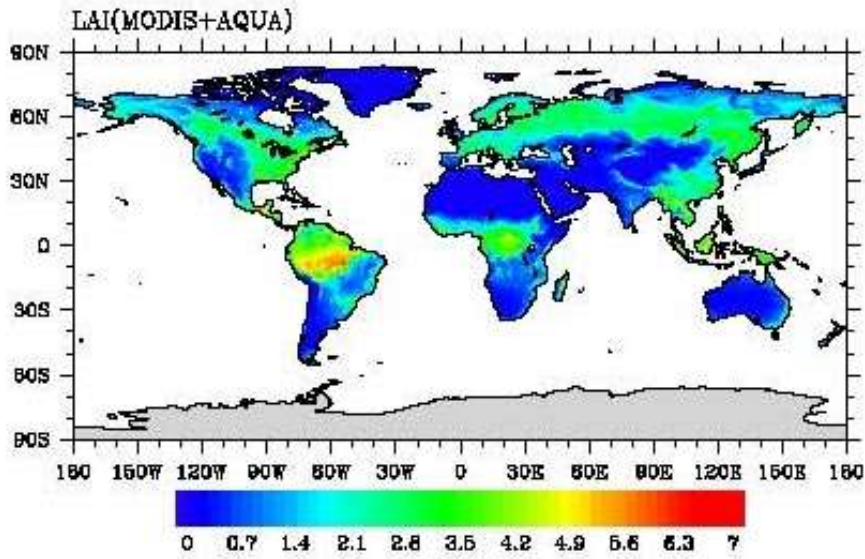
Mismatch of QLE



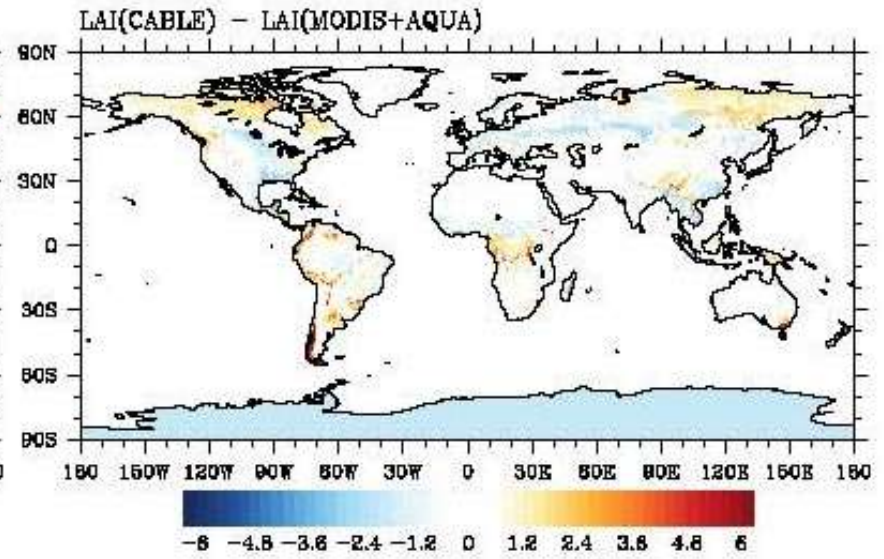
The “observed”

Mismatch

Mismatch of LAI

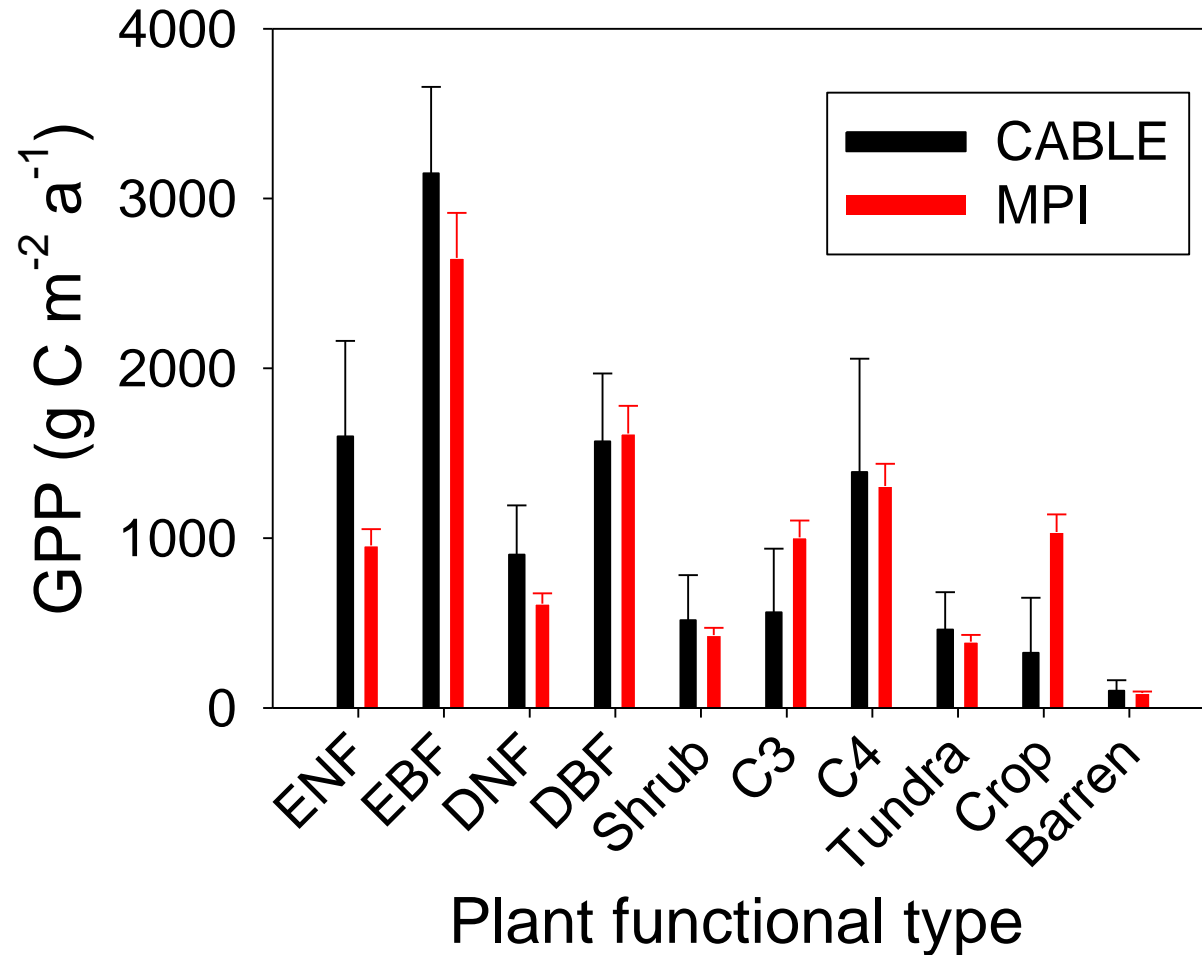


The “observed”

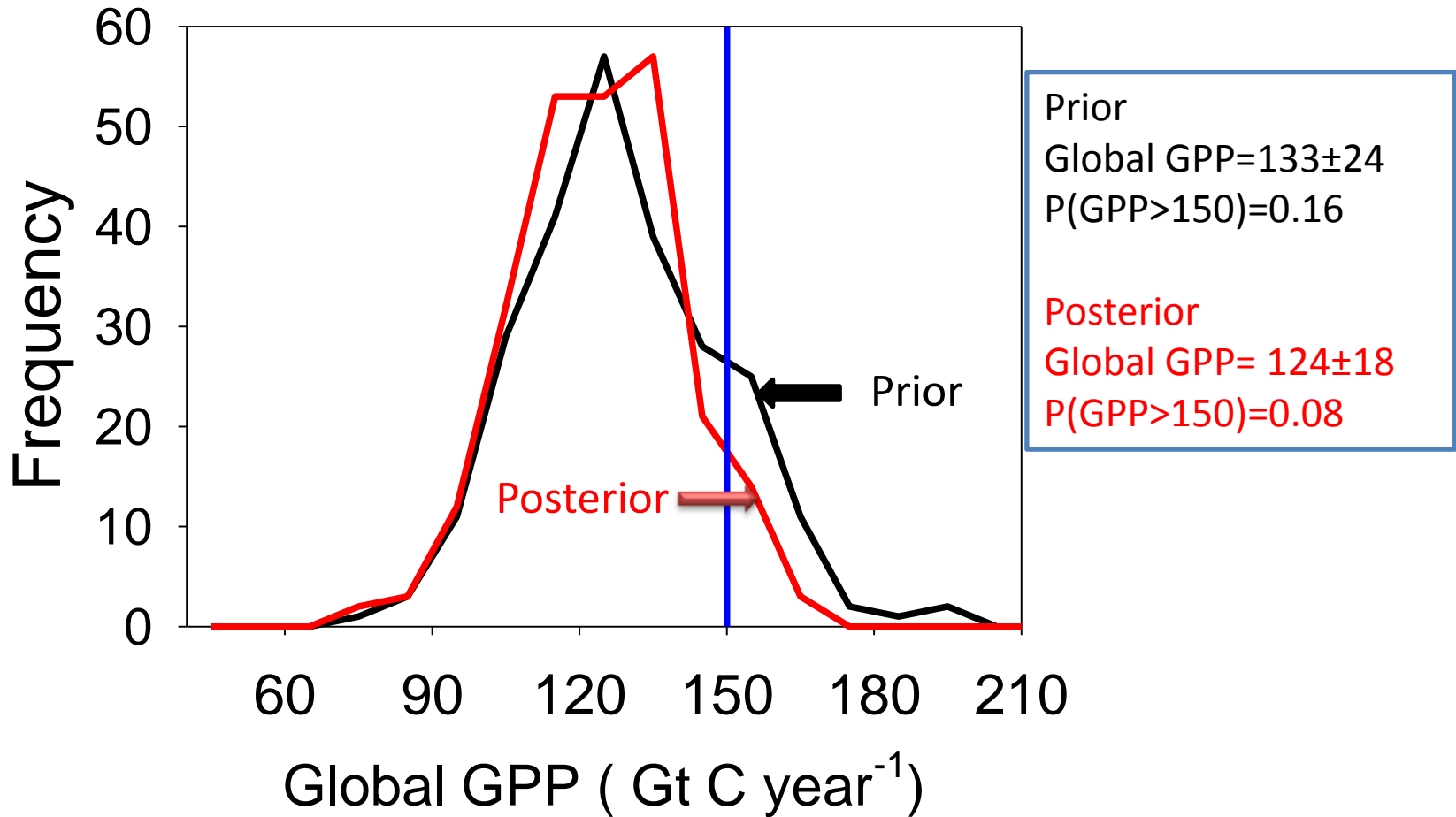


Mismatch

Predicted GPP after calibration



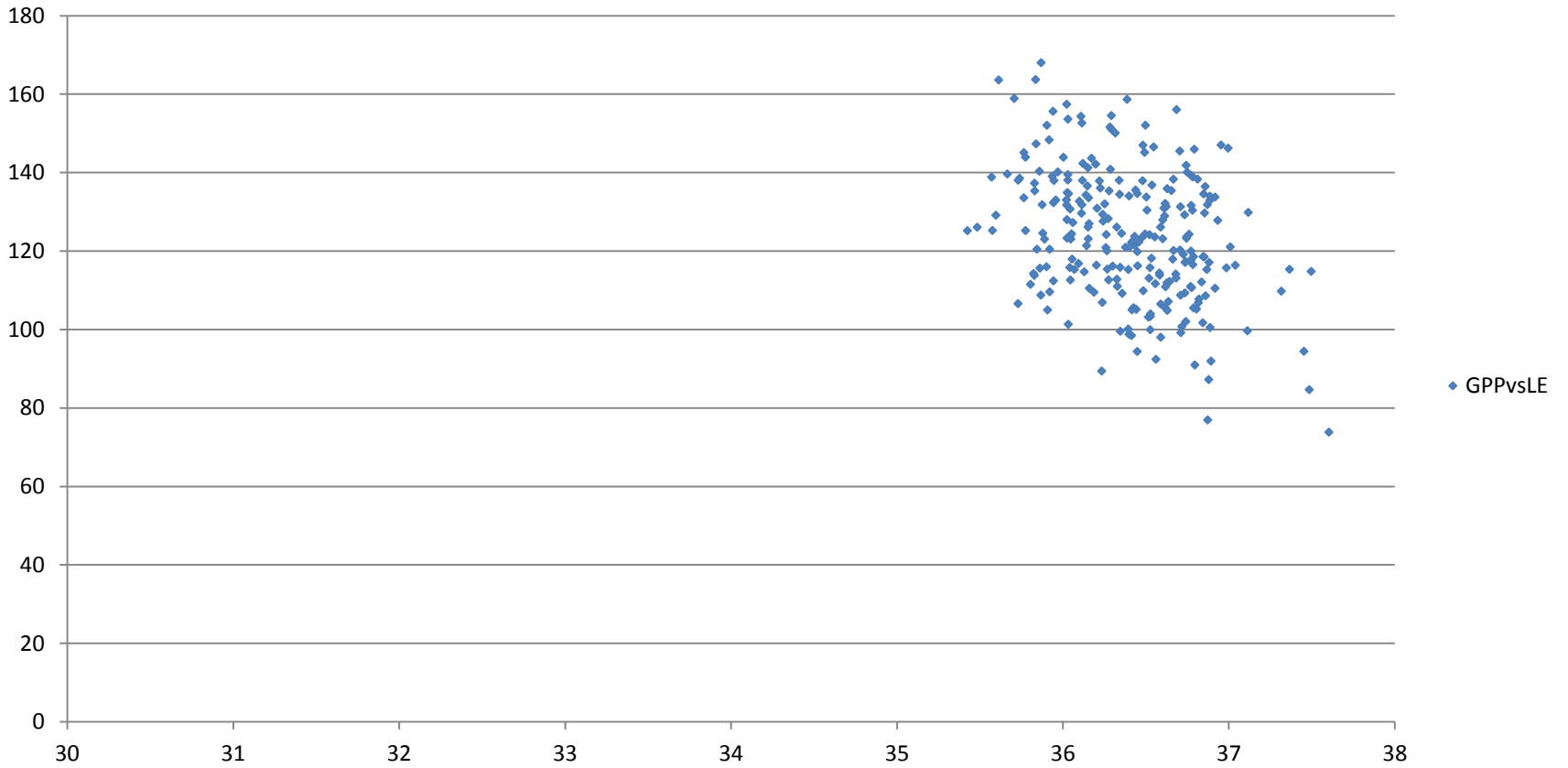
Prior and posterior GPP



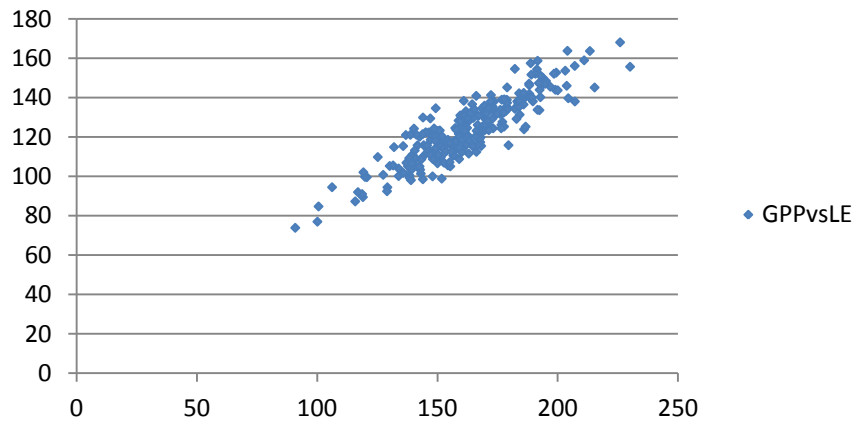
Conclusions

- Global GPP is likely to be $124 \pm \text{Gt C year}^{-1}$;
- The probability of global GPP being $>150 \text{ Gt C year}^{-1}$ is only 7%;
- Monthly LAI is a stronger constraint on global GPP than monthly latent heat flux;
- More independent estimates, such as NPP, runoff will be used to optimized more parameters and further refine the estimates of GPP.

GPP(GtC/year)vsLE(W/m2)



GPP(GtC/year)vsTVeg(kg/m2/year)



GPP(GtC/year)vsESoil(kg/m2/yea r)

