



Spatial Heterogeneity of Ecosystem Fluxes over Tropical Savanna in the Late Dry Season

Presentation by Peter Isaac, Lindsay Hutley, Jason Beringer and Lucas Cernusak

Introduction

What is the question?

- What are the main drivers of spatial heterogeneity in fluxes over Tropical Savanna?
 - ❖ Surface properties or meteorology.
 - How many leaves do we need to cover Australia?

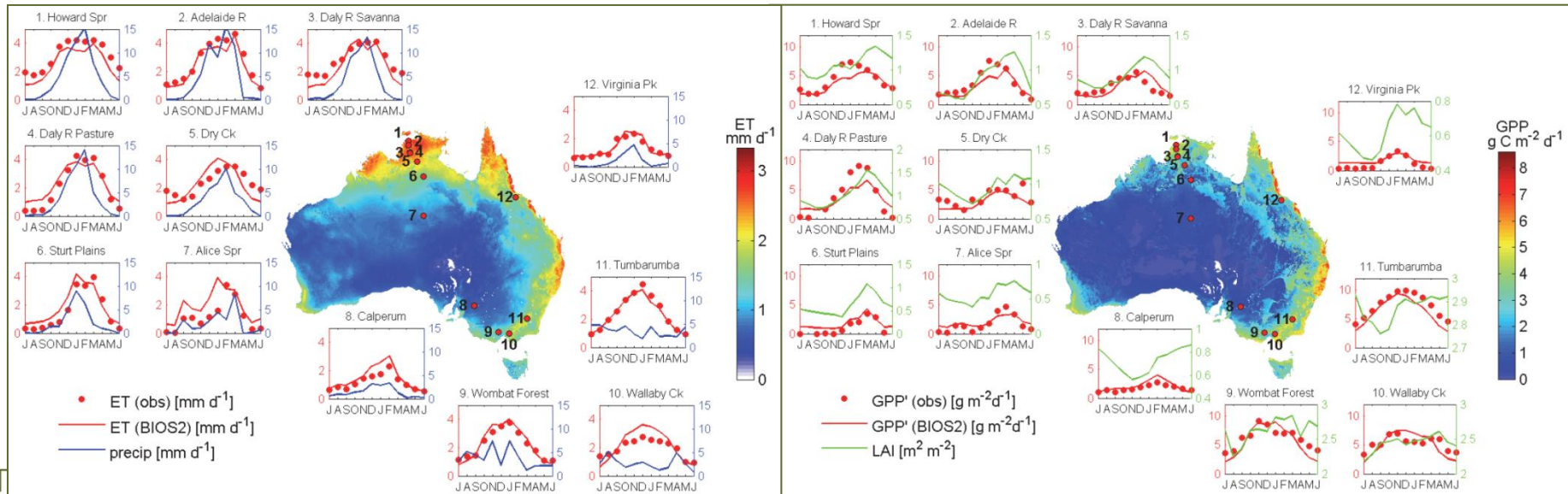
What is the data period and why?

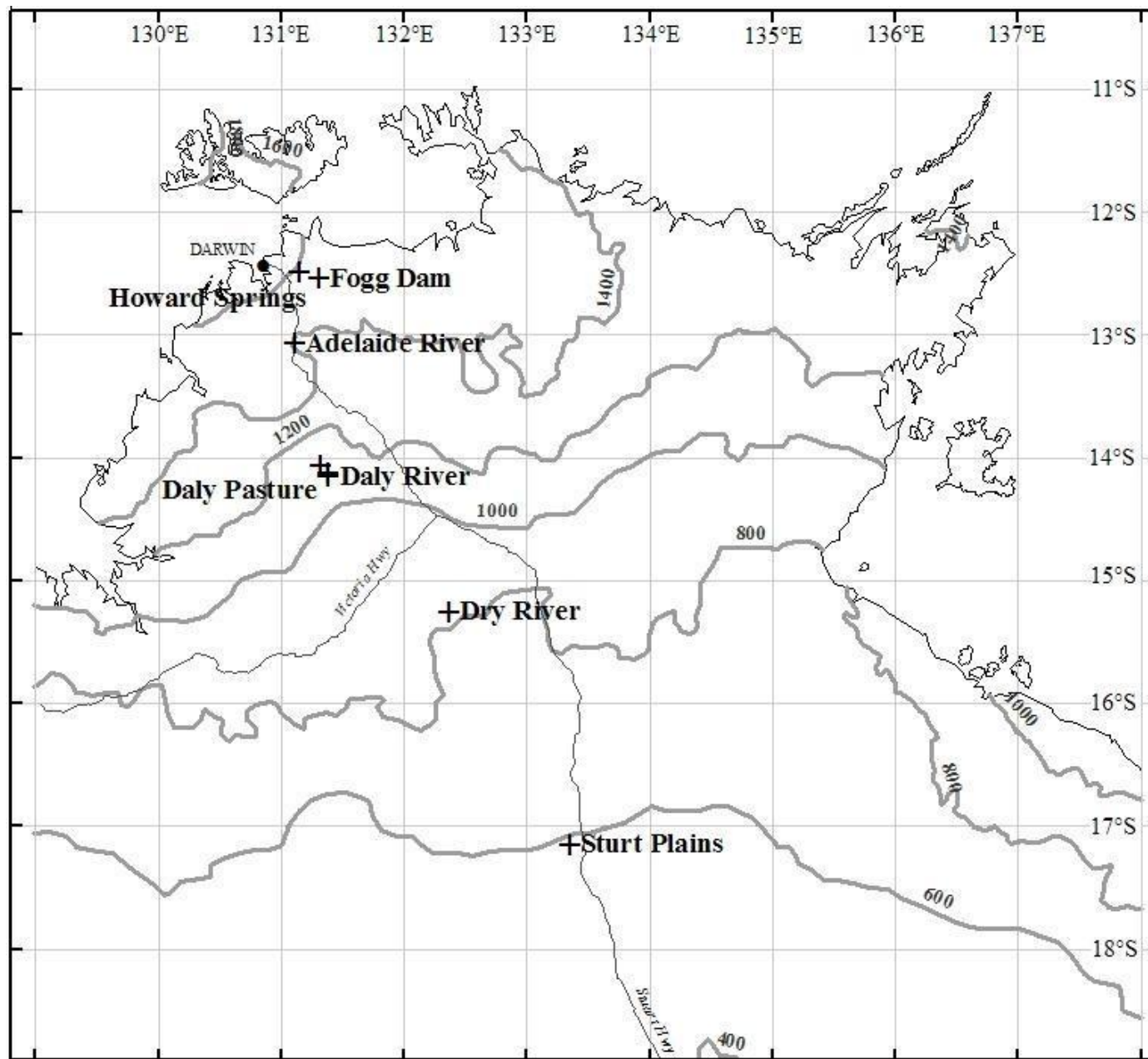
- 44 day period in the late dry season in the Tropical Savanna:
 - ❖ No C4 contribution, only C3 trees.
 - ❖ Very small soil evaporation.
 - ❖ Small, constant ecosystem respiration.

How many leaves do we need to cover Australia?

Haverd et al BIOS2 results using 6 leaves:

- “Woody” C4 sunlit and shaded.
- “Grassy” C4 sunlit and shaded.
- “Grassy” C3 sunlit and shaded.





What Could be the Drivers?

Landscape

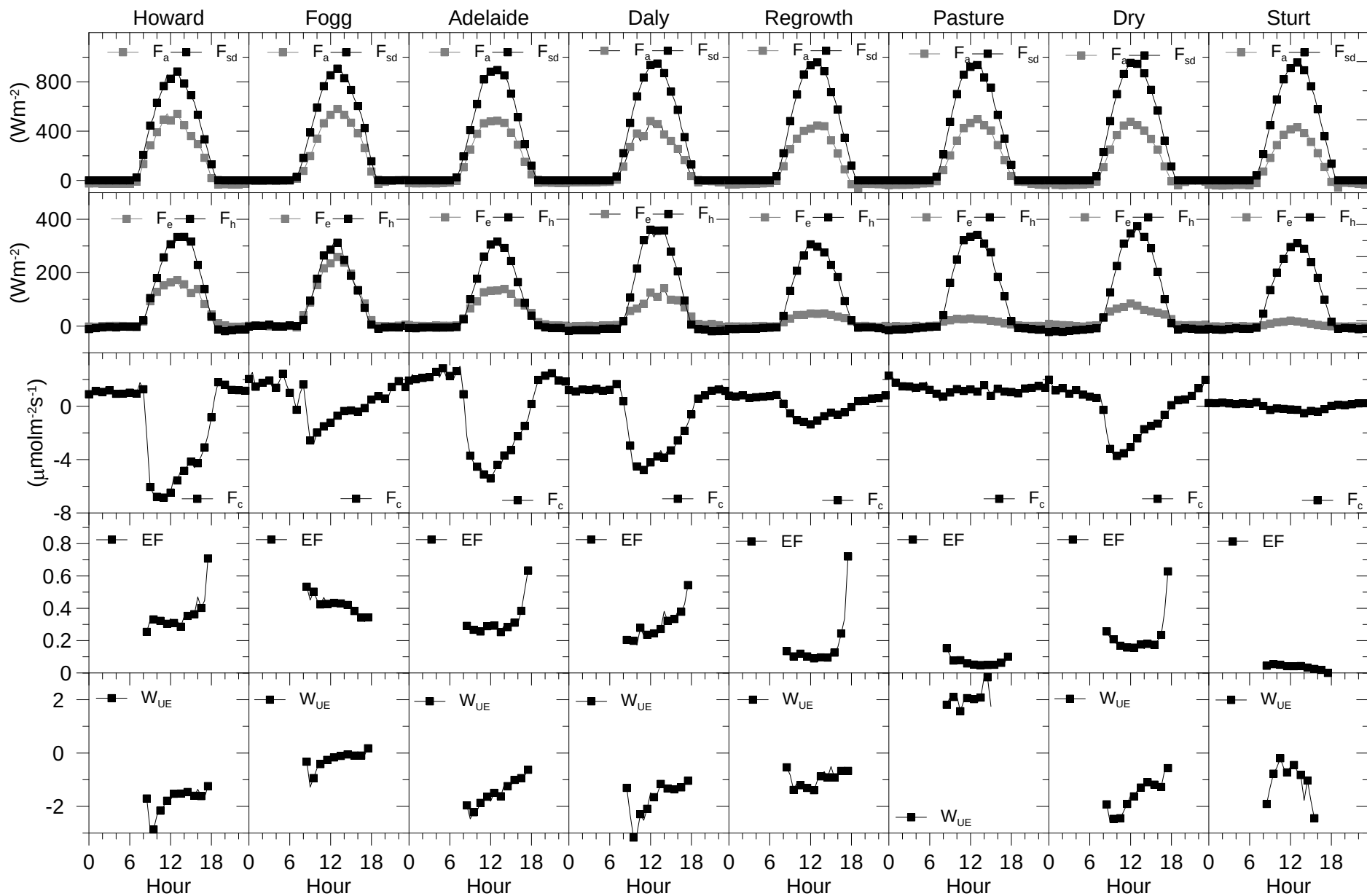
- Plant type, land use, Lai, fractional cover

Meteorology

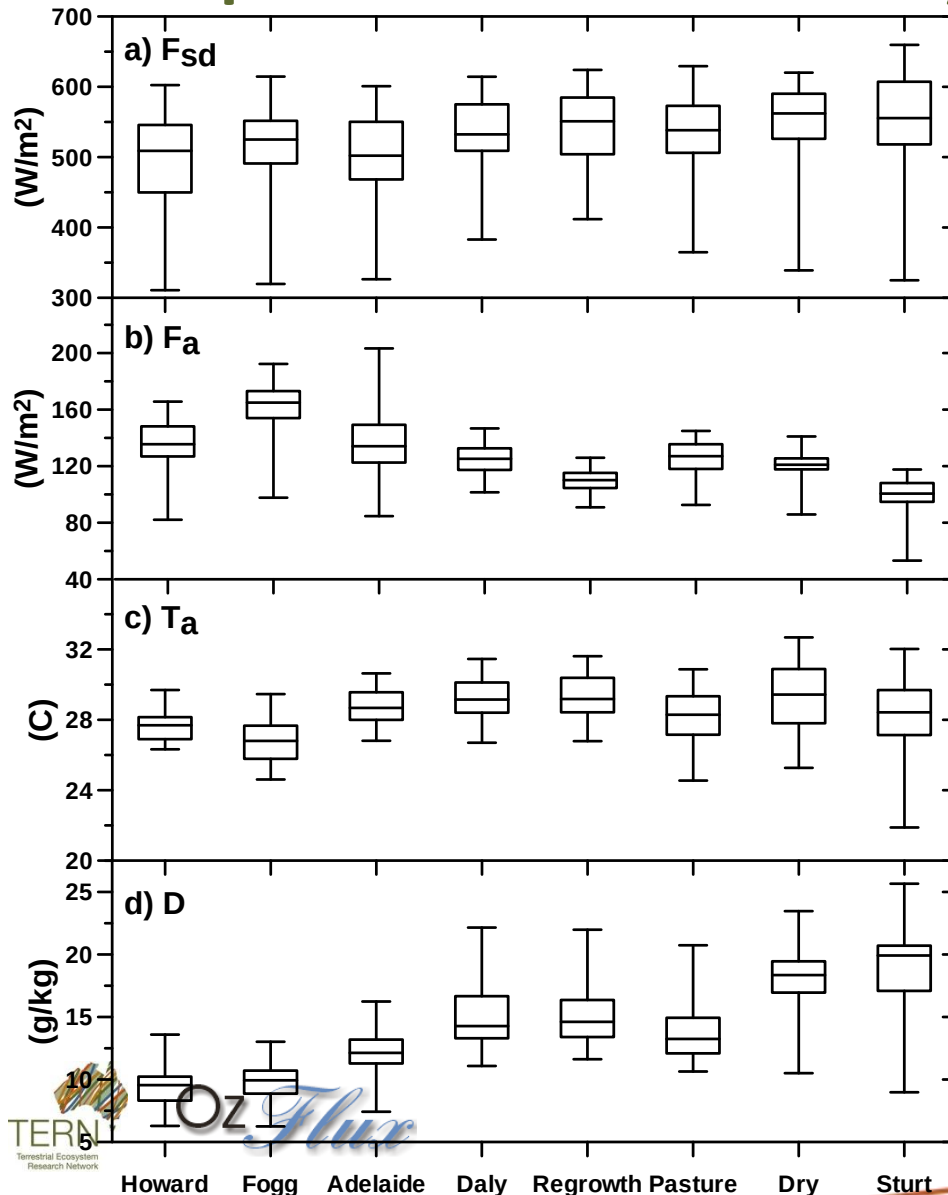
- Fsd, Fa, Ta, D, Ws, Sws,

Canopy and Leaf-level properties

- Quantum efficiency
- Maximum assimilation rate (or V_{cmax})
- Response to VPD
- Respiration



Spatial Variability along Transect



Little spatial variability in incoming shortwave, available energy and air temperature.

- Marked gradient in specific humidity deficit (D).

From the previous slide:

- Most of the spatial heterogeneity due to land use
- Concentrate on woody savanna for remainder of talk

Fluxes and Drivers along Transect

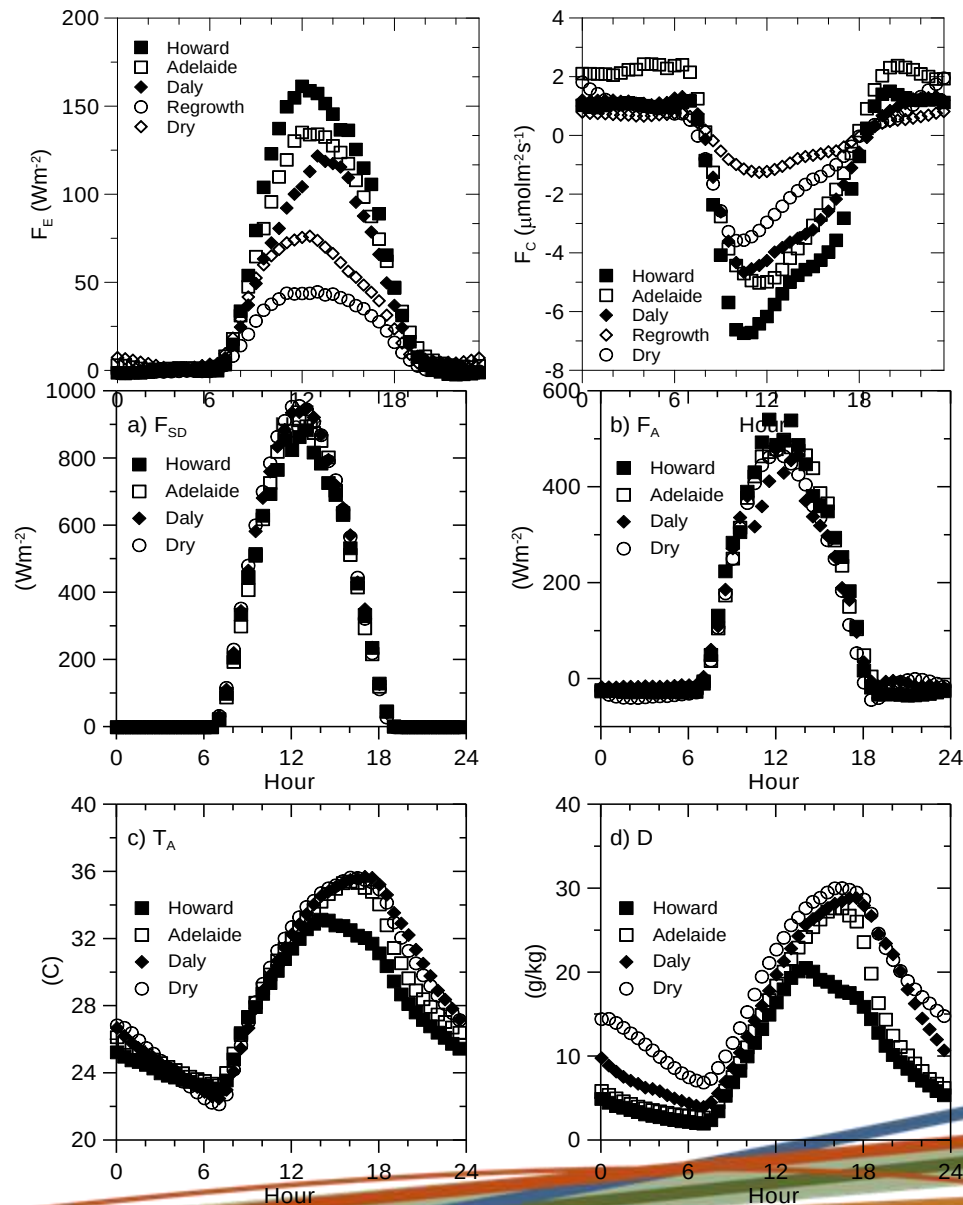
Five woody savanna sites:

- Howard Springs (1714,1.1)
- Adelaide River (1532,0.69)
- Daly Uncleared (1170,0.88)
- Daly Regrowth (1170,0.12)
- Dry River (850,0.76)

Strong gradient in fluxes.

Weak gradient in meteorology.

Daly Regrowth not used.



Comparing Response to D

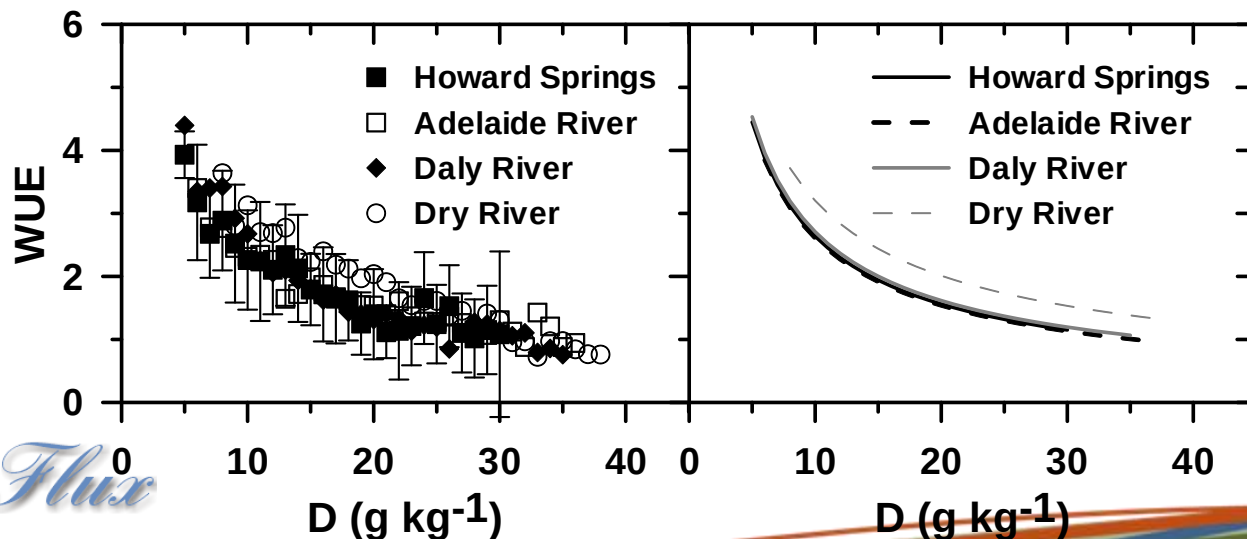
$$A = g_s (c_a - c_i) = g_s c_a (1 - c_i/c_a)$$

$$E = 1.6 g_s (q_s(T_{leaf}) - q_a)$$

$$W_{UE} = \frac{A}{E} = \frac{c_a (1 - c_i/c_a)}{1.6D} \propto \frac{1}{D}$$

$$\approx \frac{F_c - R_{eco}}{F_e} \propto \frac{1}{D}$$

Site	D exponent
Howard Springs	-0.76 ± 0.02
Adelaide River	-0.76 ± 0.02
Daly Uncleared	-0.74 ± 0.02
Dry River	-0.67 ± 0.02



Estimating R_{eco}

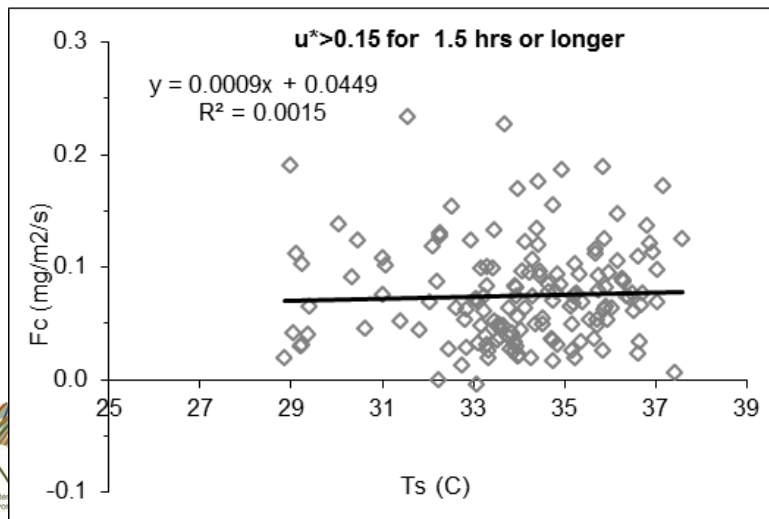
We want to partition F_c into GPP and R_{eco} so we can estimate canopy-scale quantum efficiency and A_{max} by fitting a LUE curve to day time GPP:

- Compare canopy-level parameters to leaf-level values from Lucas' work.

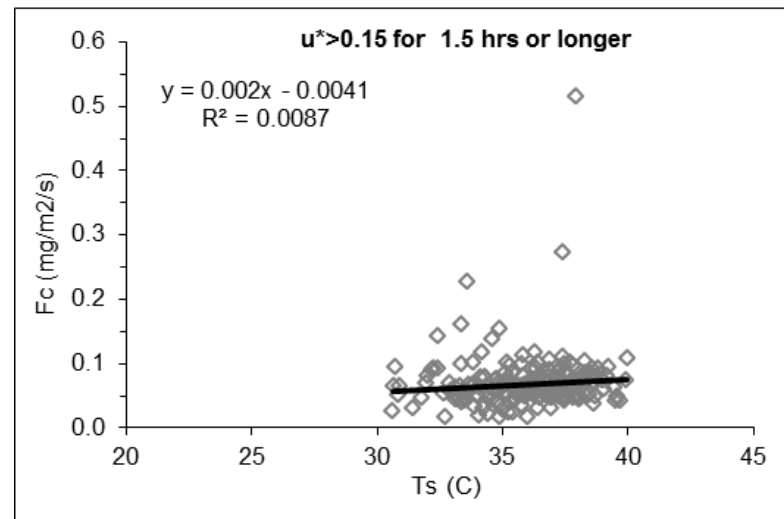
No relation found between R_{eco} with soil temperature, air temperature or soil moisture for the late dry season.

- Assumed a site-specific, constant value for 44 day period.

Howard Springs



Adelaide River



Fitting LUE Curves

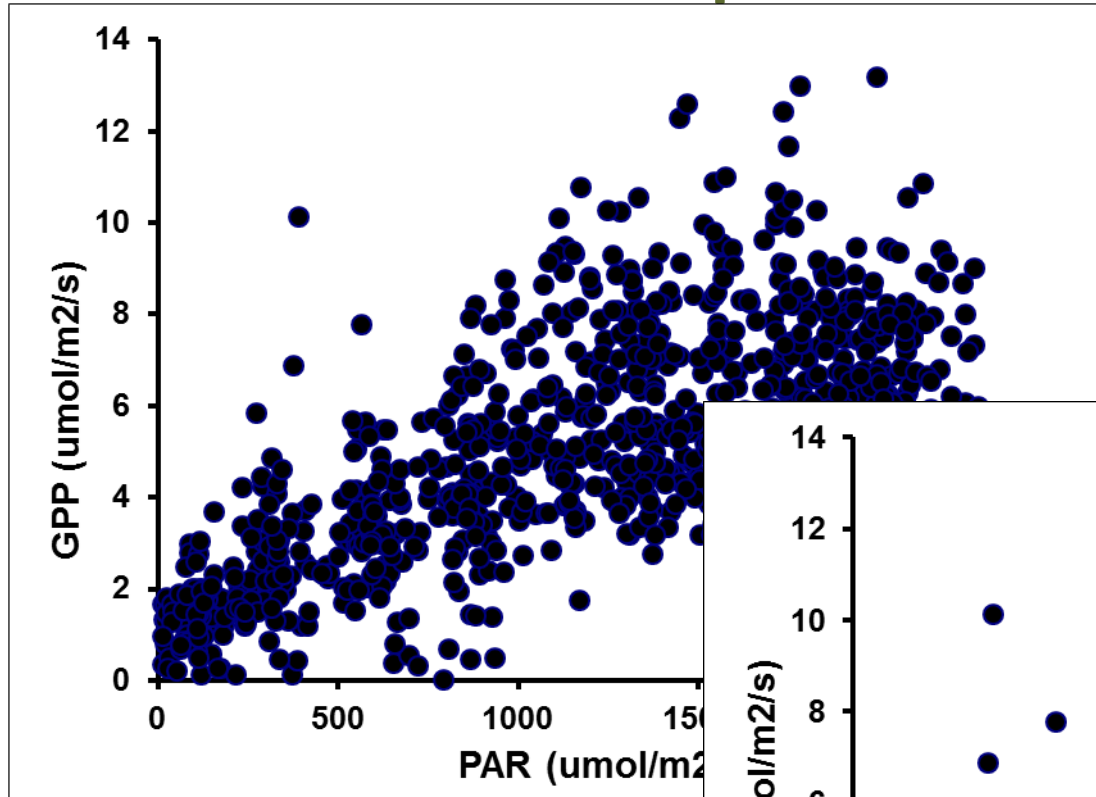
We have leaf-level photosynthetic parameters from Lucas' Li-6400 measurements:

- Quantum efficiency and $A_{\max}$
- No observable gradient in measured parameters with MAP.

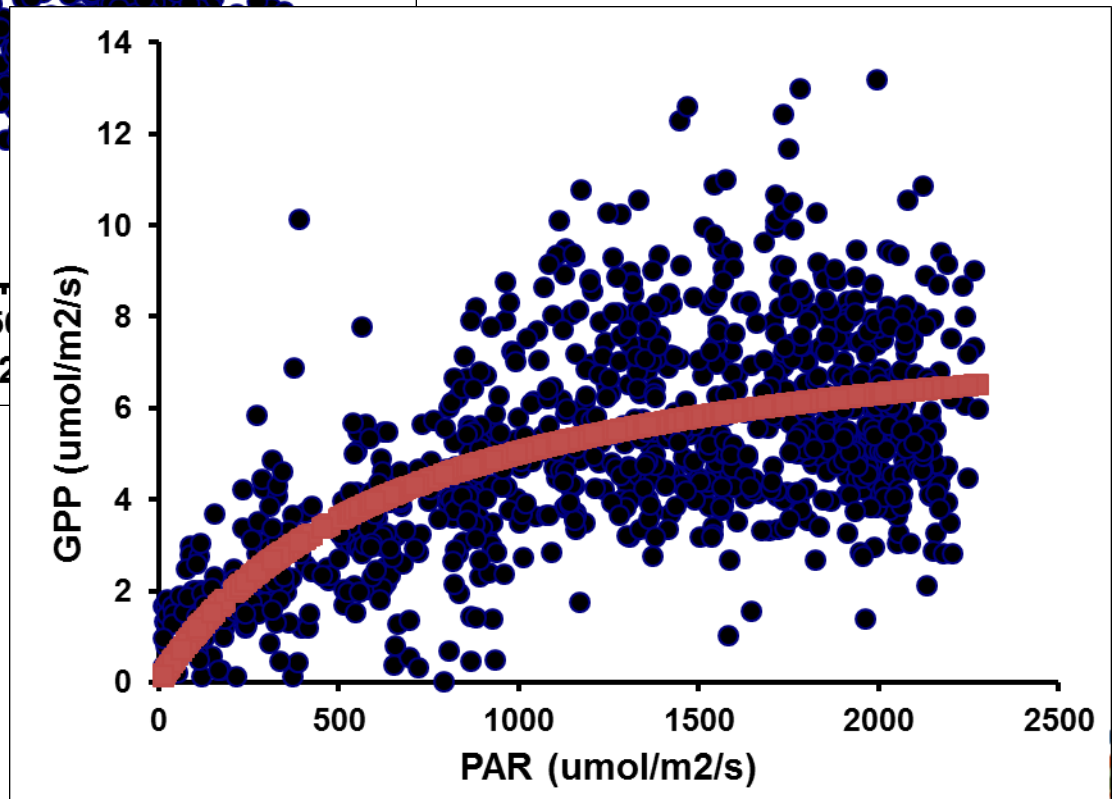
We would like to estimate the equivalent canopy-scale properties:

- Do they also show no gradient with MAP.

A Simple LUE Curve



$$GPP = \frac{\phi A_{\max} PAR}{\phi PAR + A_{\max}}$$

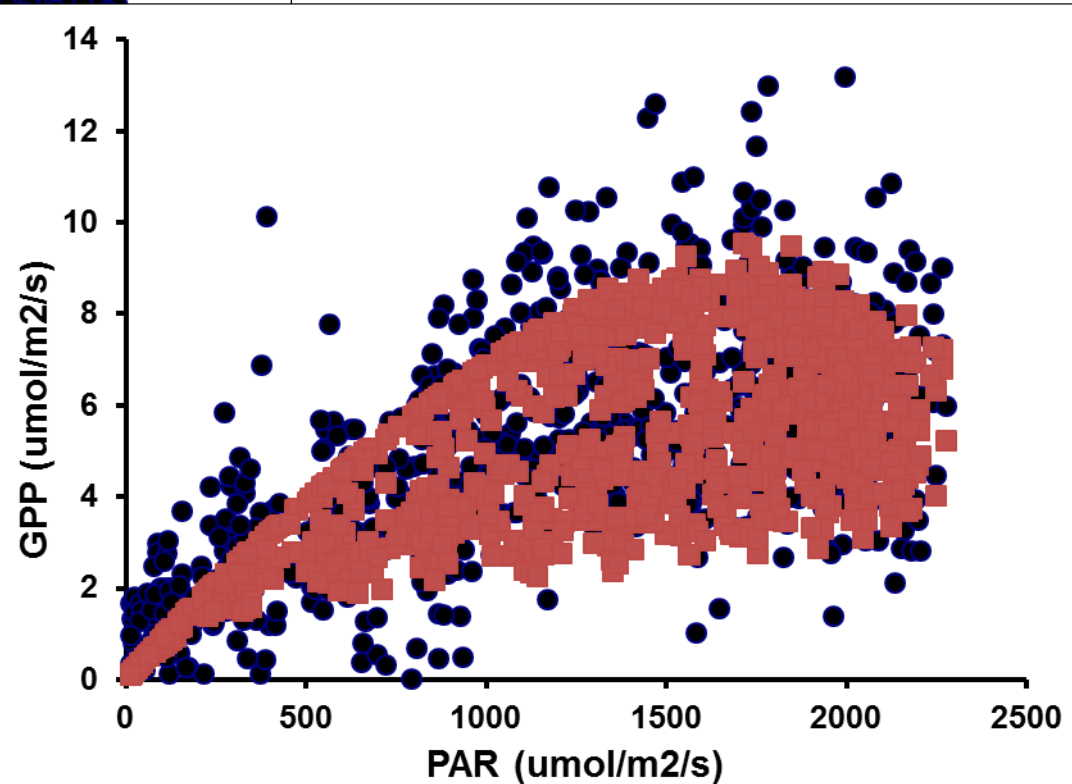
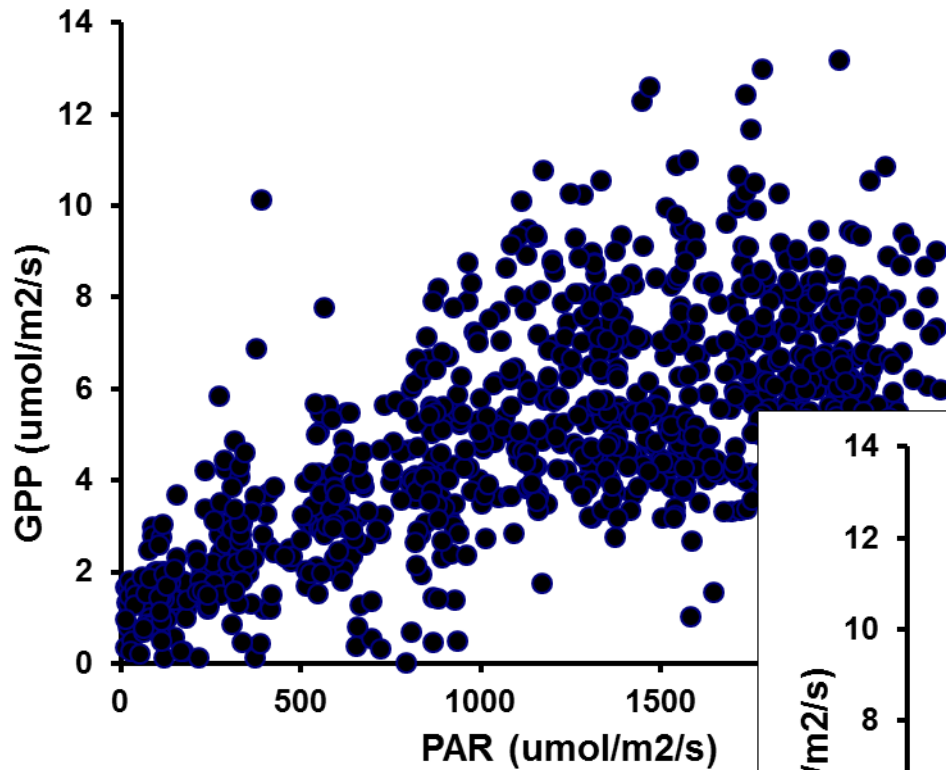


A Better Fitting LUE

$$GPP = \frac{\phi A PAR}{\phi PAR + A}$$

$$A = A_{\max}, D \leq D_0$$

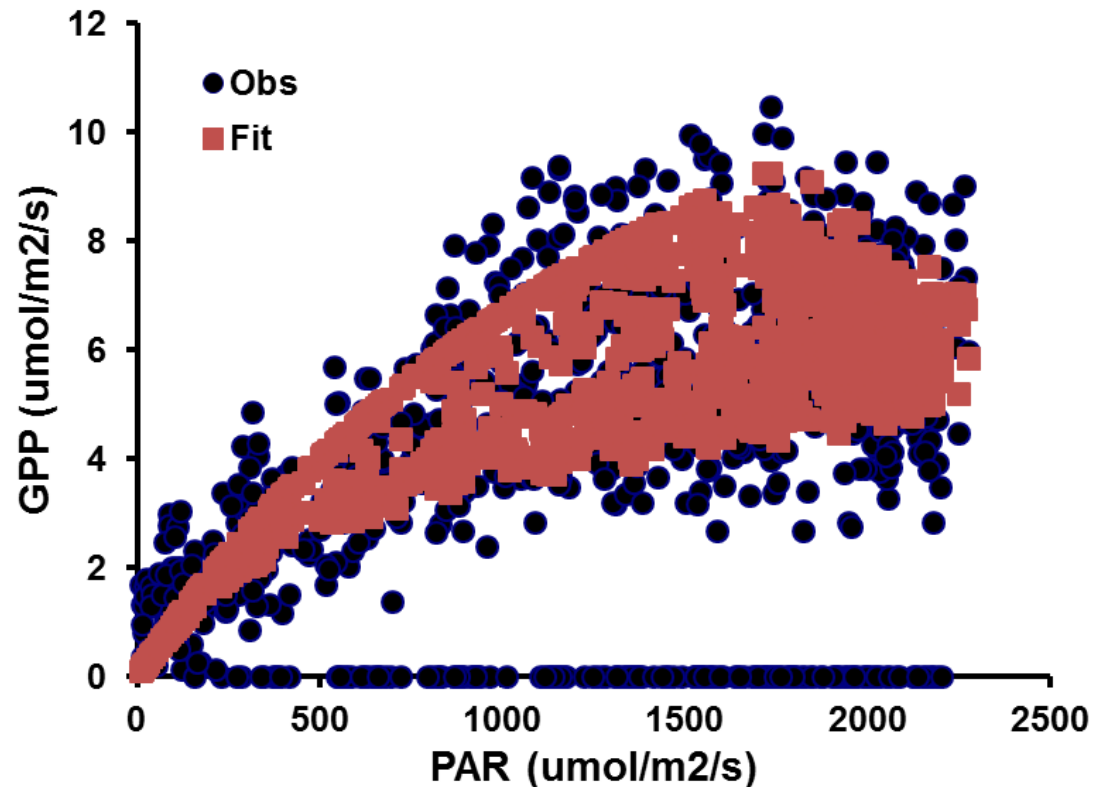
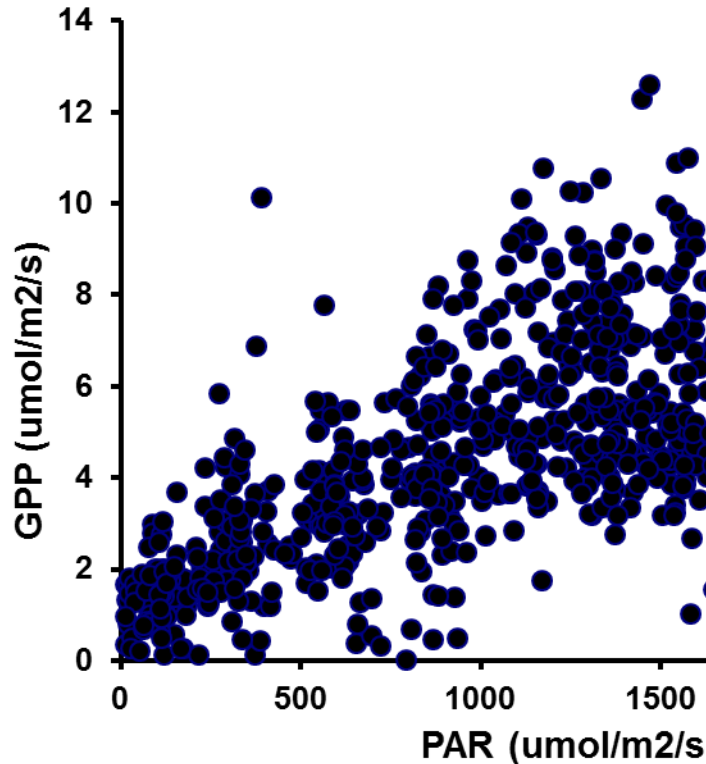
$$= \frac{A_{\max}}{\left(1 + \left(D - D_0\right) / D_0\right)}, D > D_0$$



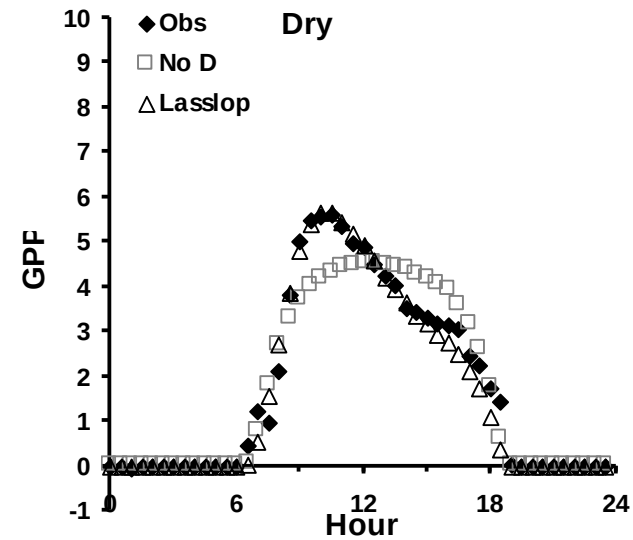
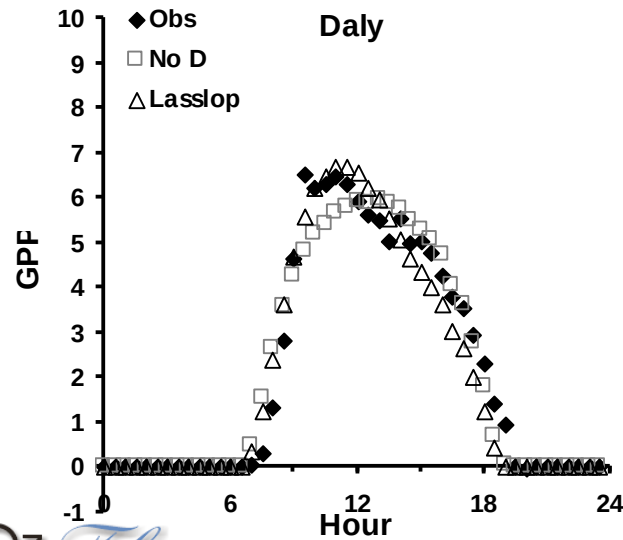
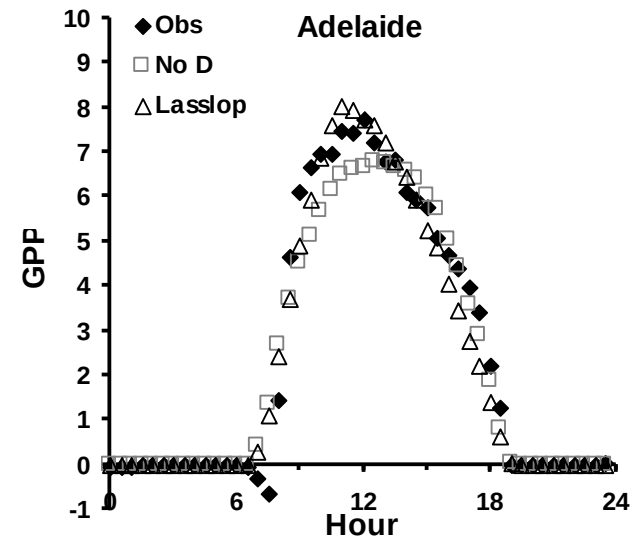
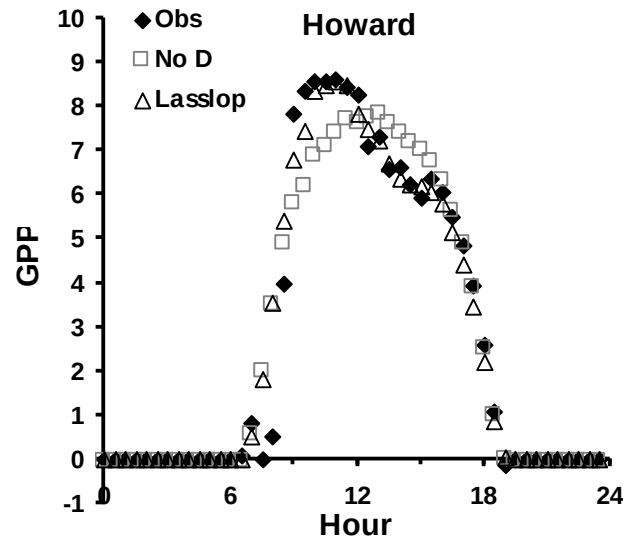
A Better Fitting Method

Robust non-linear regression

- IRLS
- Huber weight function



Diurnal Variation in GPP



Canopy Conductance and Latent Heat

We can do a similar thing with latent heat flux via a simple model for canopy conductance:

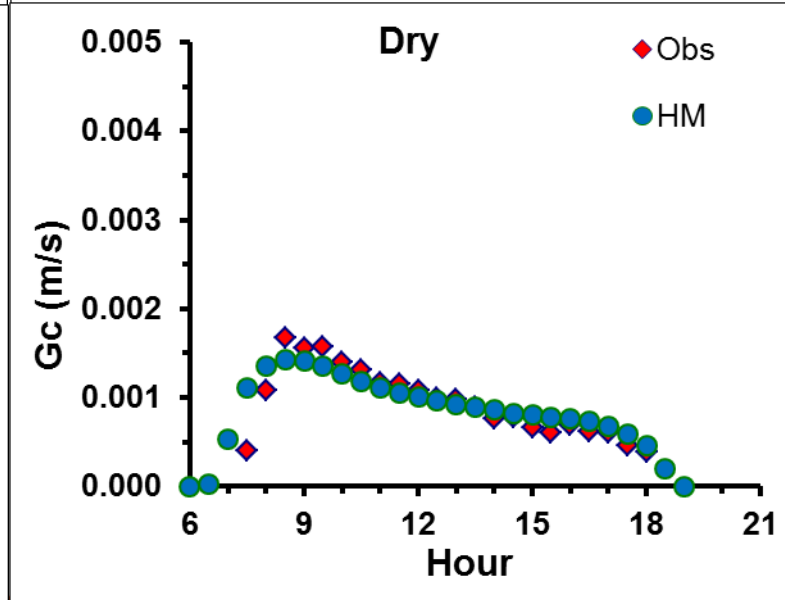
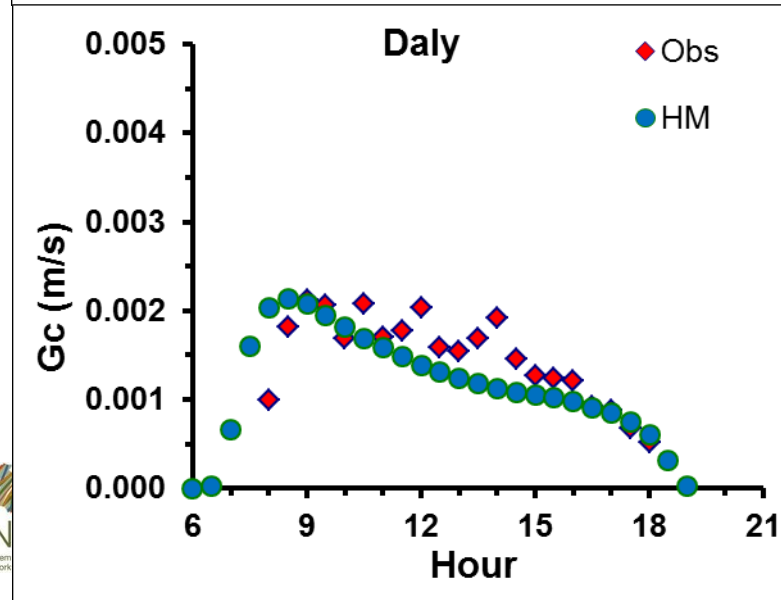
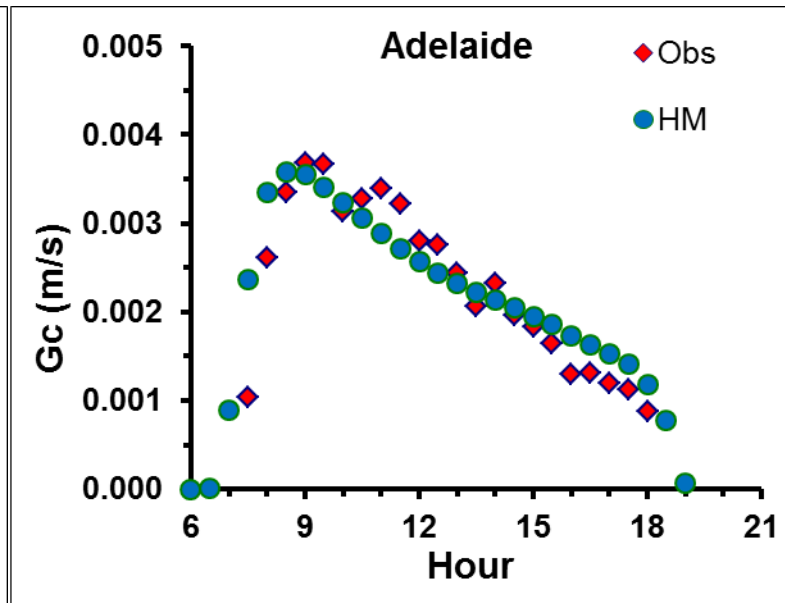
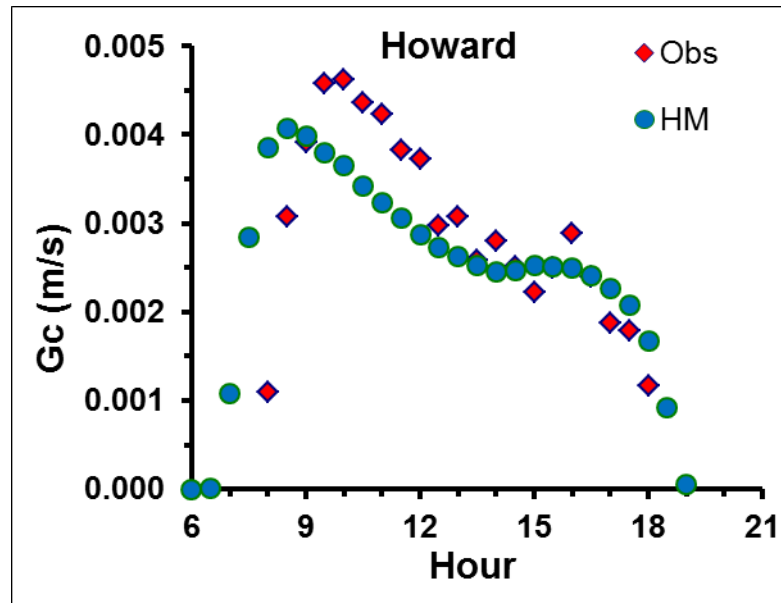
- Originally from Kelliher et al (1995), modified by Isaac et al (2004), used by Leuning et al (2008).

$$G_c = \frac{g_{sx}}{c_Q} \ln \left(\frac{S + S_{50}}{S \exp(-c_Q L_{ai}) + S_{50}} \right) f(D)$$

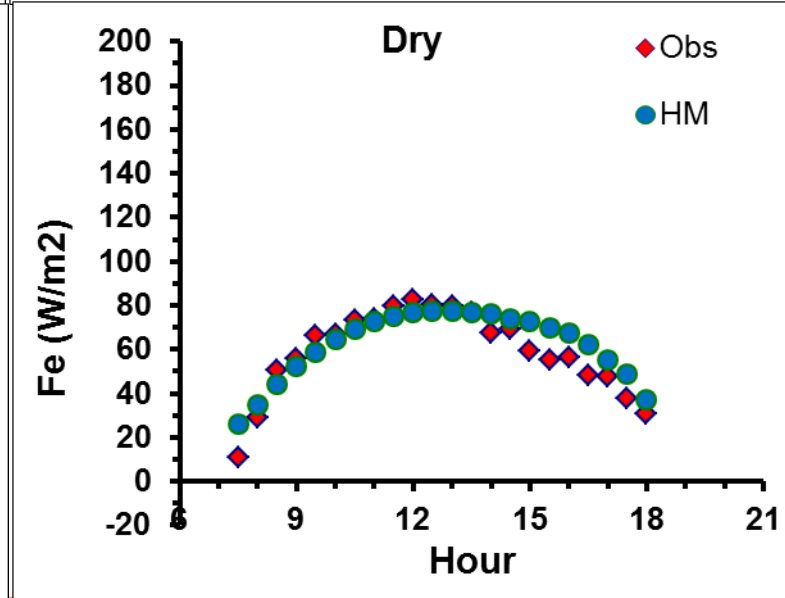
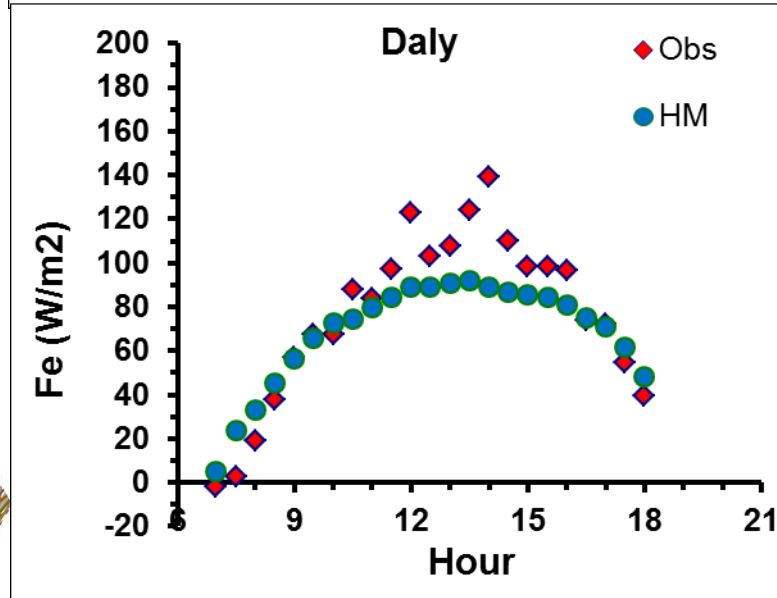
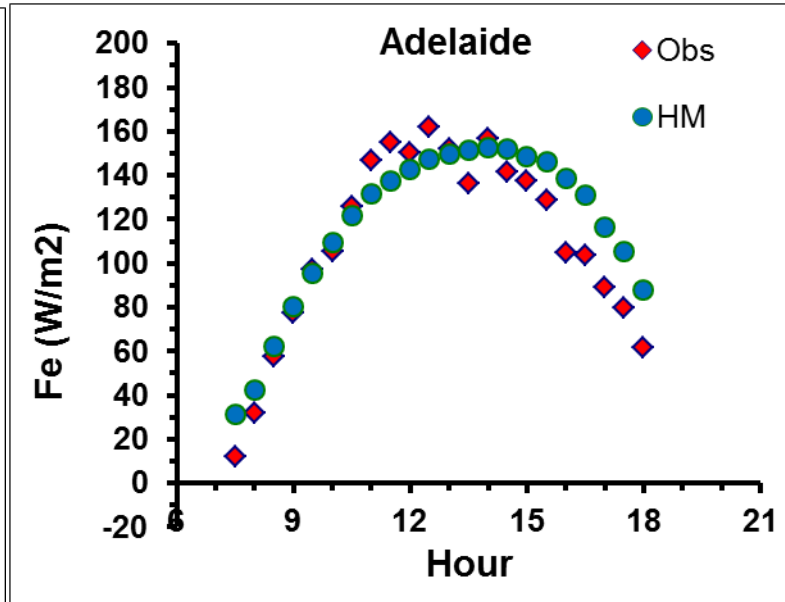
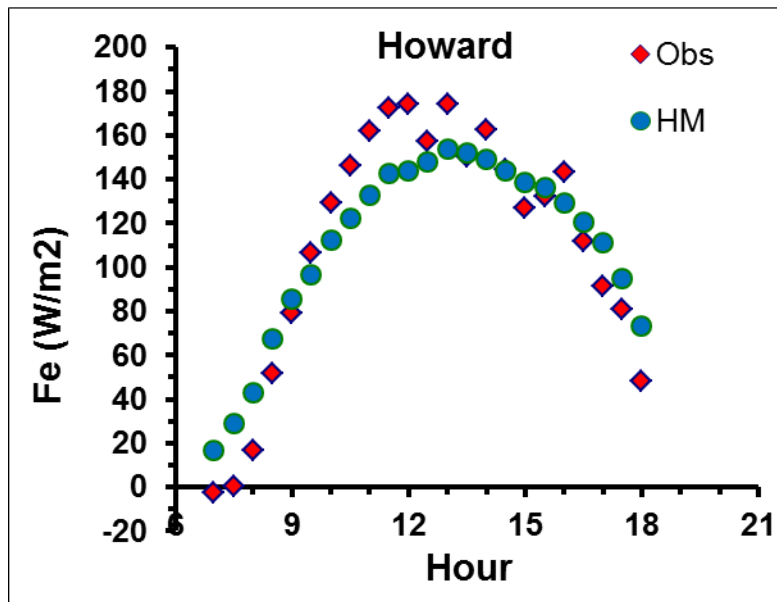
$$f(D) = 1, D \leq D_0$$

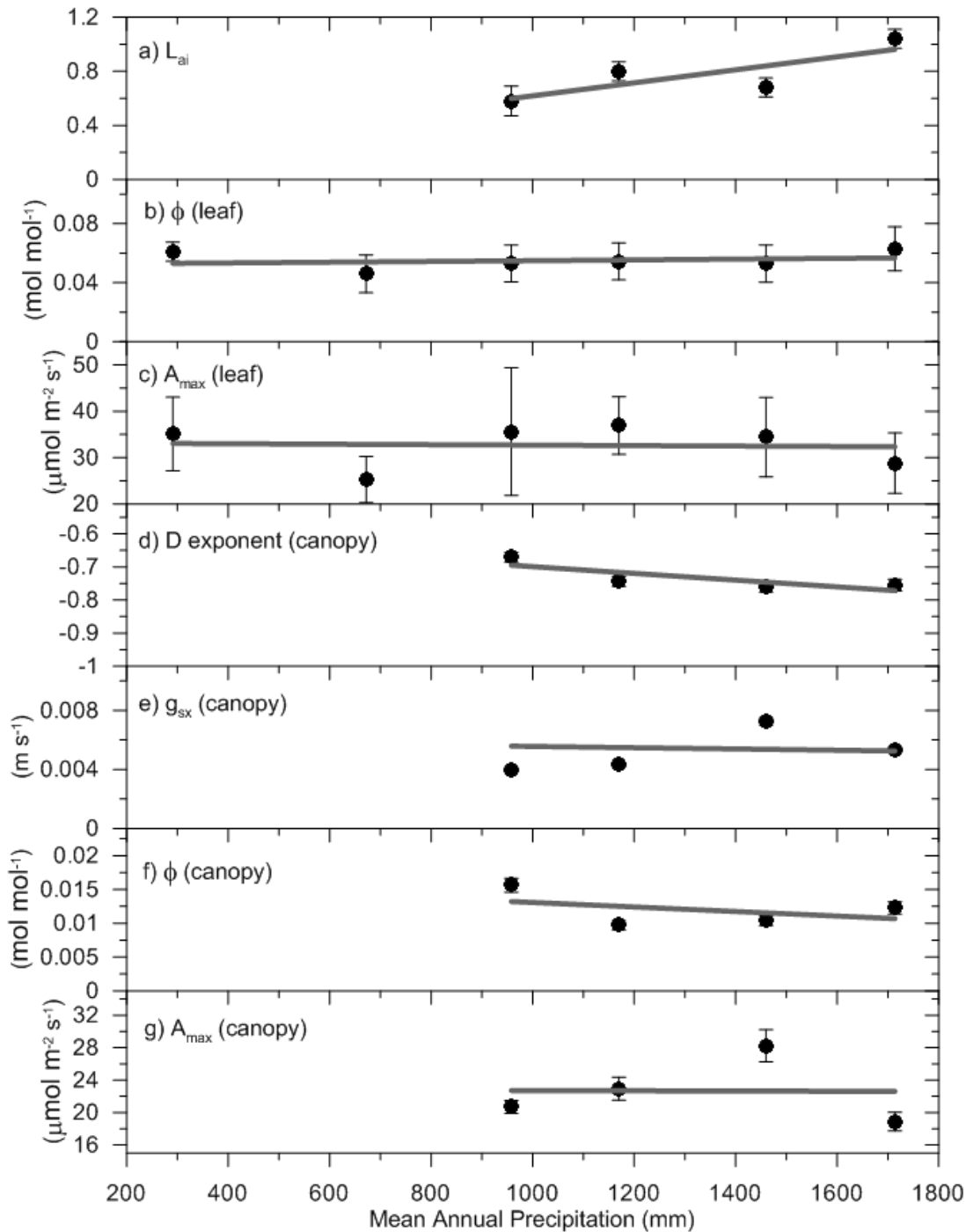
$$= \left(\frac{1}{1 + (D - D_0)/D_0} \right), D > D_0$$

Diurnal Variation in Canopy Conductance



Diurnal Variation in Latent Heat





$$m = (2.3 \times 10^{-4}, 7.3 \times 10^{-4})$$

$$m = (-3.2 \times 10^{-6}, 8.3 \times 10^{-6})$$

$$m = (-4.9 \times 10^{-3}, 3.9 \times 10^{-3})$$

$$m = (-1.5 \times 10^{-4}, -4.9 \times 10^{-5})$$

$$m = (-3.8 \times 10^{-6}, 3.0 \times 10^{-6})$$

$$m = (-8.5 \times 10^{-6}, 1.8 \times 10^{-6})$$

$$m = (-8.8 \times 10^{-3}, 8.5 \times 10^{-3})$$

Conclusions

Land use is the major driver of spatial heterogeneity over Tropical Savanna in the late dry season.

Leaf-area index is second for woody savanna:

- Photosynthetic properties don't vary with MAP for this data period.
- Response to D along transect uncertain.

We only need 1 leaf to describe the spatial variability of fluxes over Tropical Savanna in the late dry season.

Surface properties at the canopy scale may be constrained by knowledge of leaf-level properties.

Thanks

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